High Resolution Modeling of the Namibia Upwelling System during the Last Glacial Maximum

Christian Schäfer-Neth, André Paul, Michael Schulz (csn@uni-bremen.de)

DFG Research Center Ocean Margins, University of Bremen, Germany

Jan Backhaus

Center for Marine and Climate Research, University of Hamburg, Germany



Aim

Develop and run an eddy-resolving regional model to better understand the role of ocean margin upwelling for biological production and cycling of nutrients and CO_2 during extreme climatic states.



Strategy

Reconstruct sea surface conditions for the Last Glacial Maximum.

Produce regular grids from the reconstructed data.

Use the grids to drive a coarse-resolution global model.

Yield estimates for the glacial water mass distribution and the large-scale glacial circulation.

> Use the output of the glacial and modern global experiments to initialize and force the regional model.

Last Glacial Maximum SST: GLAMAP 2000 Data (°C)



 \triangle : Niebler et al. 2003. \diamond : Pflaumann et al. 2003. \Box : De Vernal et al. 2000 (Aug), Gersonde et al. 2003 (Feb). \bigcirc : Prell 1985 (Atlantic), Bigg 1994 (Mediterranean). Ice edges: Pflaumann et al. 2003; De Vernal et al. 2000; Gersonde and Zielinski 2000. Shade: SST from CLIMAP (1981).

Last Glacial Maximum SST: Gridded (°C)



Temperature and Transports in the Ventilated Thermocline (\approx 300-600 m)



- Variables stored in 1-D arrays
- Adaptive vertical coordinates
- Flooding and drying of cells in entire water column
- Open boundaries possible anywhere (e.g. on islands)
- Primitive equations with full kinematic surface and bottom boundary condition
- Implicit free surface
- Semi-implicit upstream or Arakawa-J7 advection for momentum
- Semi-implicit hybrid advection (upstream/central) for tracers
- Implicit vertical diffusion
- Variable horizontal exchange (Smagorinsky)
- Non-linear semi-implicit terrain-following seabed friction
- Terrain-following advection and diffusion at seabed
- Terrain-following isotropic resolution bottom Ekman layer







Adaptive Grid

Number of Cells in Water Column

Maximum: 66 Average: 27 \approx 300 000 wet points

Minimum: 2



- Modern: based on global experiment using SST and SSS from WOA 98
- Glacial: using global results obtained with reconstructed SST and SSS (Paul and Schäfer-Neth, *Paleoceanography*, in press)
- Annual mean forcing for 6 years to reach approximate steady state
- Initial and surface forcing T and S from global model
- Wind field interpolated from atmospheric model runs (ECHAM/T42, G. Lohmann and S. Lorenz)
- SST/SSS restoring time constant: 30 days
- Open boundary forcing by T, S, and surface elevation from global model



Results: SST ($^{\circ}$ C)



• The upwelling region shifts from $28^{\circ}S$ to $24^{\circ}S$ in the glacial run

Results: Glacial minus Modern Temperature Anomaly (°C)



- The anomaly increases at the surface, but decreases at depth
- These opposing trends can be explained in terms of changed upwelling

Results: Temperature along 27° S, 0–500 m depth (°C)



• The lower boundary of coastal upwelling rises to 150 m (modern) and 250 m (glacial)

- In the glacial experiment, the upwelling becomes more intense and broader
- In the modern experiment, the cooling at depth is more pronounced

Results: Horizontal Velocity at 250 m depth (m/s)



- Standing eddies are generated in spite of smooth annual wind fields
- Topographic features dominate influence of T/S forcing
- Currents are slightly stronger in the glacial experiment

Results: Vertical Velocity at 150 m depth (m/d)



- Both experiments develop quite similar patterns
- Coastal upwelling intensifies in the glacial experiment
- Highest velocities are found over the ridge and the seamounts
- Coastal upwelling reaches only 10-50% of open ocean upwelling

- In the regional model, topographic influence on hydrography and circulation appears to be much stronger than surface forcing
- Temperature and depth of upwelled water are realistically represented in the regional model
- Location and intensity of coastal upwelling change in accordance with the glacial SST reconstructions
- Include glacial-interglacial topography changes
- Force with seasonal T, S, and surface elevation
- Include daily wind variability
- Employ heat and freshwater fluxes instead of T/S restoring
- Develop biological sub-model: N, Si, functional plankton groups, Fe, Dust