The 1993 Northeast Water Expedition Data Report on CTD Measurements of RV "Polarstern" Cruises ARKTIS IX/2 and 3

Die Nordostwasser-Polynia Expedition 1993 Datenband über die CTD-Messungen während der Expeditionen ARKTIS IX/2 und 3 mit FS "Polarstern"

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Introduction

This data report includes all recorded downcast CTD-profiles of Polarstern cruise ARKTIS IX, leg 2 and 3. It corresponds to two other volumes of Berichte für Polarforschung, i.e.

The 1993 Northeast Water Expedition Scientific cruise report of RV "Polarstern" Arctic cruises ARK IX/2 and 3, USCG "Polar Sea" cruise NEWP and the NEWLand expedition. Edited by H.J. Hirche and G. Kattner, Ber. Polarforsch. 142 (1994)

and

The 1993 Northeast Water Expedition

Data Report of RV "Polarstern" Arctic Cruises ARK IX/2 and 3. Edited by G. Kattner and H.J. Hirche, Ber. Polarforsch. 145 (1994). This report adds a graphical listing of parameters measured with the CTDsystem. In addition, a short summary of main results is given as a first

orientation to water masses in the area. Bottle sample data can be found in Kattner and Hirche (1994).

The two legs of RV "Polarstern" cruise ARKTIS IX were part of the International Arctic Polynya Programme (IAPP) and lasted from May 16 to June 24 (leg 2), 1993, and June 26 to August 4, 1993 (leg 3).

Instrumentation and data processing

CTD casts have been performed with a Seabird 911+ system, equipped with a standard sensor set and a number of additional optical sensors. The standard set consists of a pumped temperature-conductivity-system to optimize time alignment and time constants of the respective sensors. Optical sensors include measurements of chlorophyll flourescence, backscattering, yellow substance fluorescence (all three Dr. Haardt instruments), light intensity relative to the surface (LICOR with AWI interface) and transmission (Seatech instrument). Yellow substance measurements have not been recorded on stations 105 to 123.

Four calibration checks have been performed in the central Greenland Sea (75°N, 3°W) at about 3000 m depth on May 23, June 22 and 27, and August 2, since owing to the high spatial gradients the NEW area is not suited to calibrate CTD sensors by comparative temperature measurements and water sampling.

Evaluation of the calibration casts yields an accuracy of better than 5 mK for temperature and 0.006 for salinity during the entire duration of the cruise. The data have been processed by standard methods using mainly Seabird software. This includes optimizing time alignment between conductivity and temperature measurements, optimizing their time response, correcting for the thermal mass of the conductivity cell, suppressing of possible upward CTD movements during downcasts, and application of a median filter over 5 scans for the fluorescence sensors.

Summary of main results

In this summary the water masses on the northeast Greenland shelf are described with respect to their relationship to the surrounding waters in Fram Strait, Greenland Sea and Polar Arctic Ocean in order to give an introduction to the hydrography of the area. More detailed information can be found in the papers listed as references at the end of this text.

The bathymetry of the northeast Greenland shelf is dominated by a trough system - consisting of Belgica Trough, Norske Trough and Westwind Trough - around the shallow area of Belgica Bank (Fig. 1). The troughs are 300 to 500 m deep, Belgica Bank is partly shallower than 40 m. A persistent fast ice feature (Norske Ø Ice Barrier) bridges over Norske Trough at about 79°N and the combined effect of this mechanical barrier together with a northward coastal current gives rise to a summer polynya - the Northeast Water (NEW) - downstream of the ice barrier (Schneider and Budéus, 1994; Schneider and Budéus, 1995). A second shallow area, Ob Bank, is found at the northern boundary of the northeast Greenland Shelf.

Earlier researchers already observed a generally two-layered structure in the area (Aagaard and Coachman, 1968; Bourke et al., 1987). The top layer is occupied by fresh (S about 32.5) and cold (T close to freezing temperature) waters categorized as Polar Water (PW) and assumed to be imported by the East Greenland Current (EGC), while the the fact that deeper waters in the troughs are warmer (up to 1°C) and more saline (up to 34.9) has been attributed to the influence of Return Atlantic Water (RAW). In contrast to these ideas the NEW expedition revealed that deeper waters on the shelf are clearly different from RAW, and that the fresher water on the shelf is formed locally and should be distinguished from PW.



Fig. 1. Bathymetry and topographic features of the Northeast Water area (from Bourke et al., 1987). Henrik-Krøyer-Islands are abbreviated HKI. In the area covered by fast ice, depth contours are dashed.

Polar Water / East Greenland Shelf Water

The import of PW from the north into the NEW area can be excluded due to three main arguments. First, the northward flowing coastal current (North East Greenland Coastal Current) inhibits a direct inflow. This northward current has been observed by a number of researchers (Kiilerich, 1945; Bourke et al., 1987; Schneider and Budéus, 1994; Schneider and Budéus, 1995) and shows geostrophic speeds of about 0.1 m/s. It turns to the east over Westwind Trough and is confined to the upper water column. It gives rise to an anticyclonic surface circulation north of 79°N.

Secondly, the temperature structures of the upper water column are different in the NEW area and north of Ob Bank. In the NEW a layer which stays close to freezing temperature is maintained throughout the summer. This layer is found in about 70 m depth at salinities of roughly 32.5. At this salinity waters north of Ob Bank show their maximum deviation from freezing temperature. Even if this water would be imported into the NEW, it would have to be cooled in the NEW area during a winter season, and thus be formed by a local formation process.

A further argument can be found when regarding nutrient concentrations, in particular silicate. The very high silicate values from north of Ob Bank do not enter the NEW area (Budéus et al., 1996).

Consequently, this water body should be distinguished from PW stemming from the Arctic Polar Ocean, and it is termed East Greenland Shelf Water due to its local origin.

Intermediate Waters

Paquette et al. (1985) introduced a third water body, which is named Knee Water (KW) after its shape in a TS-diagram. This water seems to be of Atlantic origin, cooled and freshened in the Arctic under the contact of ice (Moore and Wallace, 1988). This water can serve excellently as a tracer for the circulation on the shelf.

From a mixing triangle in the TS-space the maximum concentrations of KW observed in 1 m bins of each CTD profile can be calculated (Budéus and Schneider, 1995). It can be seen that this water spreads diagonally over the shelf but does not enter the NEW with the NEGCC. This observation is further supported by the distribution of nutrient concentrations and yellow substance fluorescence and corroborates the local character of the shelf system.

Deeper Waters

As indicated by their temperatures, deeper waters in the trough system around Belgica Bank cannot be formed locally by convection. There is much evidence that vertical mixing during winter - whether convection or wind mixing includes only the uppermost 100 to 150 m. The high temperatures of the deeper waters indicate an ultimate origin from Atlantic waters, but a direct inflow of RAW can be excluded. There is an abrupt change from the high temperatures of RAW to lower maximum temperatures in the troughs, and the temperature maximum there shows virtually no horizontal gradient (Budéus and Schneider, 1995). This opposes the idea of a continuous supply and dilution of RAW. Further evidence is given to this by nutrient data, which show much higher silicate concentrations in the troughs (8 to 12 µMol/I) than in RAW (about 6 μ Mol/l) (Budéus et al., 1996). From a transect along the troughs' axes it can further be seen that there is no one-directional flushing of the trough system but deeper waters do enter from the mouths of both Belgica Trough and Westwind Trough. This indicates a long residence time of these waters, what is confirmed by Helium/Titium ages of 14 to 20 years (R. Bayer, pers. comm.).

Conclusion

The northeast Greenland shelf seems to be occupied by waters of mainly local character. The properties of the upper water column are determined mainly during winter by cooling and vertical mixing. To maintain roughly similar salinities in the East Greenland Shelf Water, the fresh water input from local land runoff in summer has to be balanced on longer terms by salt input from brine release or entrainment of saltier waters from below.

The deeper waters in the troughs show some resemblance to waters observed north of Ob Bank and in the Polar Arctic Ocean, but presently there is no evidence for a clear relation to these waters. Deep water formation by vertical convection and brine release can be excluded with respect to the Greenland Sea.

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Data

The locations of stationwork during the two legs of ARKTIS IX are shown in Fig. 2 and 3. A station list with exact positions can be found in Hirche and Kattner (1994), Ber. Polarforsch. 142, p. 167 ff.

Naming convention

The naming convention for stations and casts is as follows: The profile is described by 8 characters; the first 4 characters indicate the cruise (ark9), the next 3 characters indicate the station; the last character stands for the cast number, starting with zero for the first cast of a station. Thus ark91230 is the first profile at station 123.

Comments on profiles

The profiles are depicted for the upper 500 m since most participants are interested in the uppermost few hundred metres only. Deeper casts are cut off to resolve the upper 500 m adequately. The first 10 to 20 m may contain apparent instabilities due to the ship's manouvers on stations (see e.g. ark90470). No attempt has been made to correct for this. Optical parameters, in particular chlorophyll, occasionally show profiles with spikes in spatial scales of true variabilities. These can be identified by observing simultaneously the transmission signal, which has also to decrease at locations of high chlorophyll concentrations. The entire data set has been plotted with the same scales to allow an easy comparison between different profiles. As a consequence some extreme salinity and density values are out of range. Although we agree that the resulting plots do not look exactly perfect, clipping has been suppressed for those instances to still provide some information then.



Fig. 2. Stations during ARKTIS IX/2



Fig. 3. Stations during ARKTIS IX/3











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