A18 FAMOS Workshop Woods Hole Oceanographic Institution November 1-4, 2016

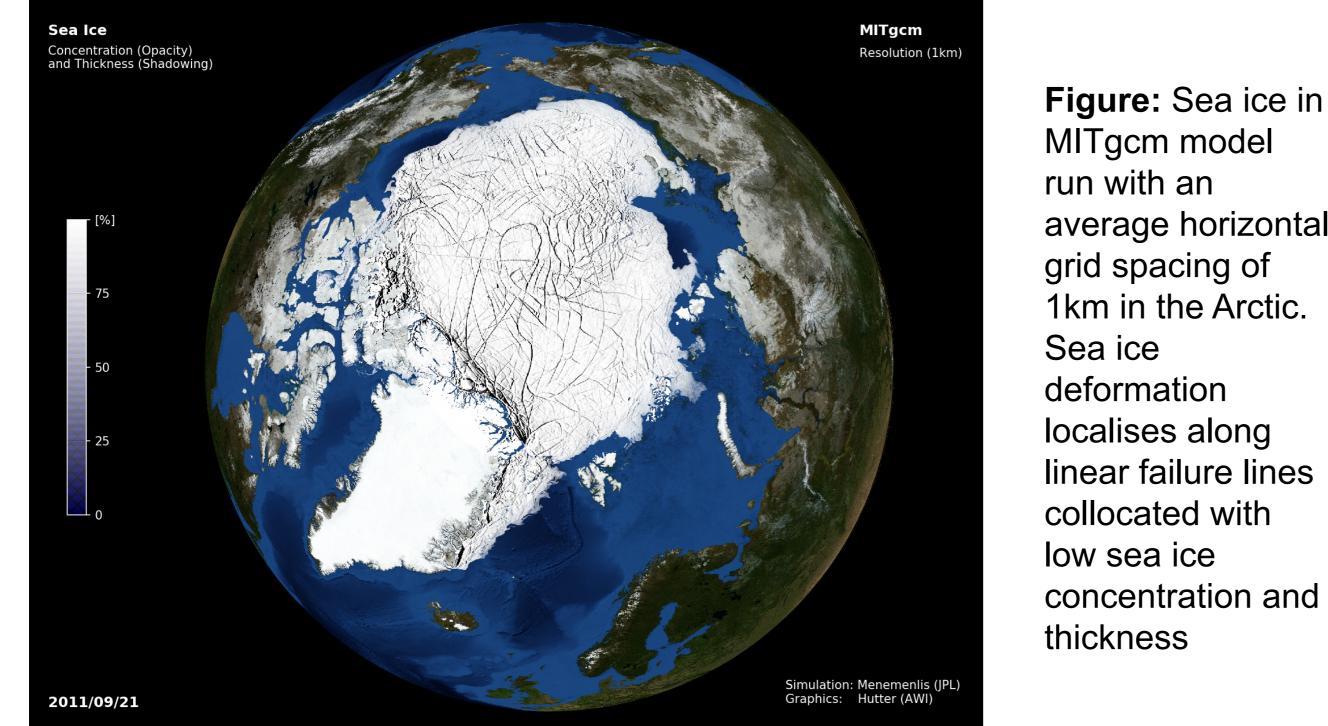


Nils Hutter¹ (Image in the image is a construction of the im

Scaling properties of Arctic sea ice deformation in high-resolution viscous-plastic sea-ice models

Leads in viscous-plastic (VP) models

VP sea ice models at coarse resolution are known to reproduce statistical and scaling properties of sea ice deformation inappropriately [Girard et al., 2009], but ...



Conclusions

- The resolved leads improve strongly the scaling properties of deformation rates in VP models.
- Arctic wide model analysis shows agreement with other RGPS studies and experiments with the EB-rheology.
- Seasonal and regional variation of spatial scaling is captured by the model.

MITgcm model run with an average horizontal grid spacing of 1km in the Arctic. Sea ice deformation localises along linear failure lines collocated with low sea ice concentration and thickness

➡ At very high resolution leads emerge in viscous-plastic sea ice models.

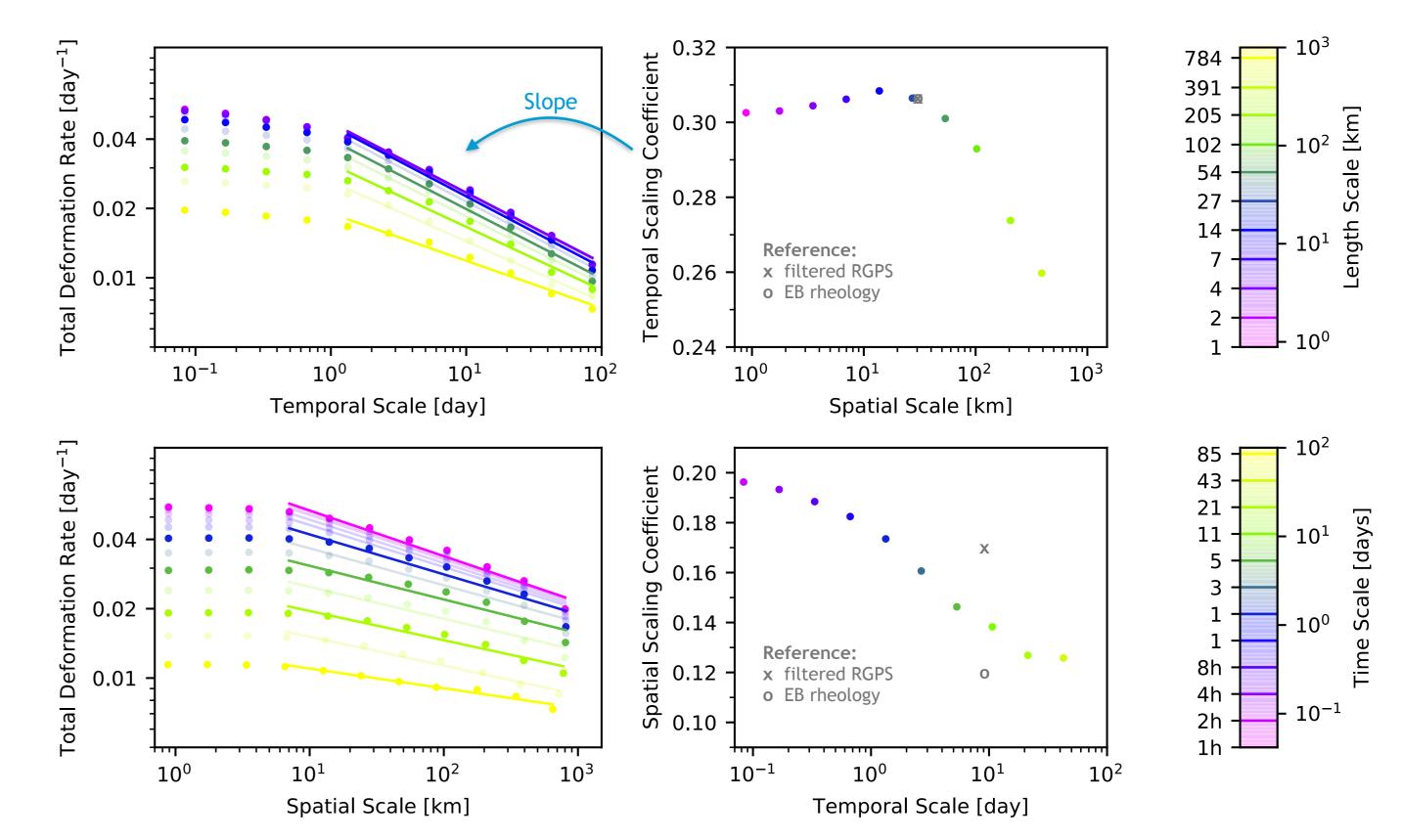
Research Objectives: Do the emerging leads in VP sea ice models at very high resolution result in scaling properties of sea ice deformation comparable to satellite observations?

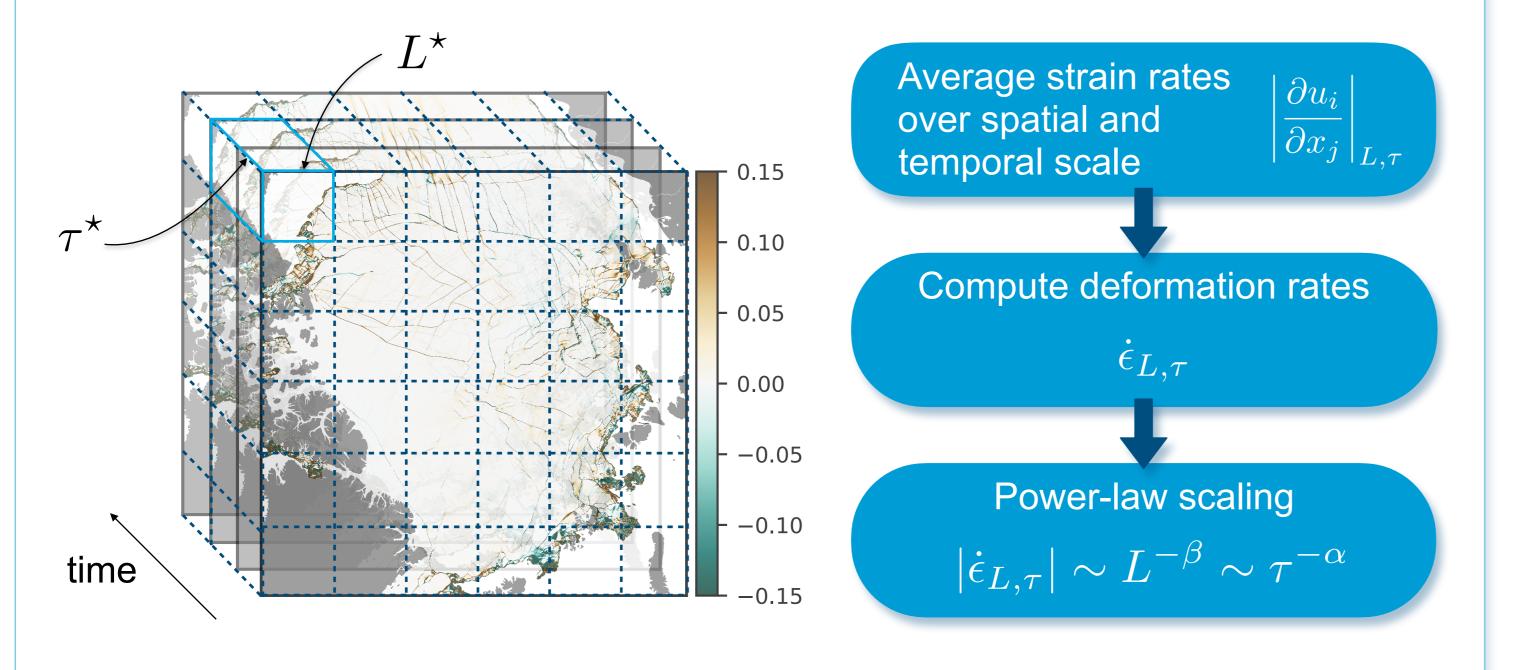
Scaling analysis

Adaption of the scaling analysis of *Marsan et al., 2004* for velocities on regular grids with integrated temporal scaling analysis.

- A more comprehensive model evaluation requires satellite data with larger spatial coverage along with higher temporal resolution and longer model simulation.
- VP rheology appears to be an appropriate framework for modelling sea ice deformation at high resolution

Modeled scaling properties





Low scaling exponent —— similarity across scale —— propagation of stress High scaling exponent — heterogeneity across scale — local deformation events

Evaluation with satellite data

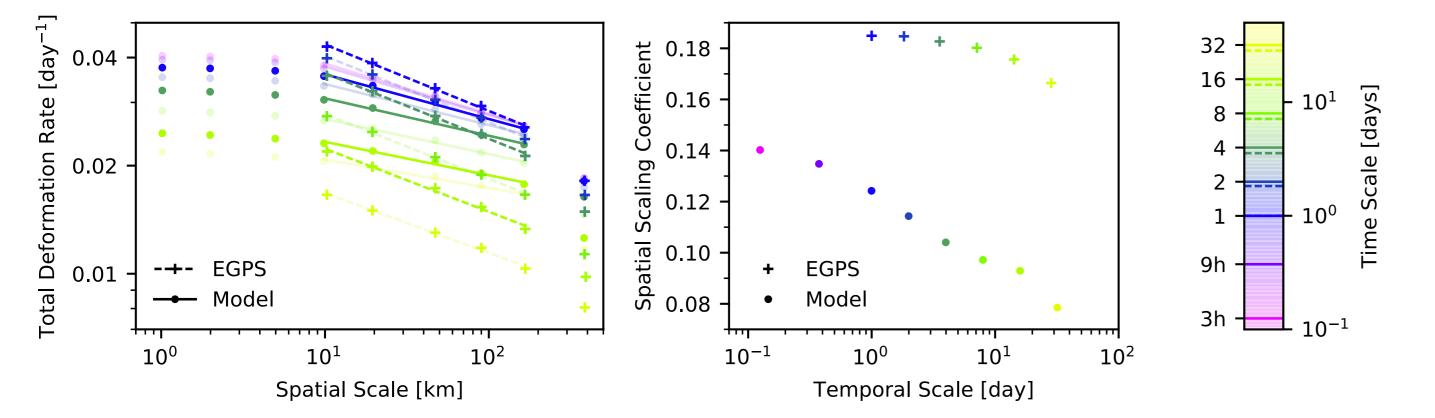


Figure: Temporal and spatial scaling properties of sea ice deformation (left column). Temporal and spatial scaling are coupled (right column). Reference data from Rampal et al., 2016

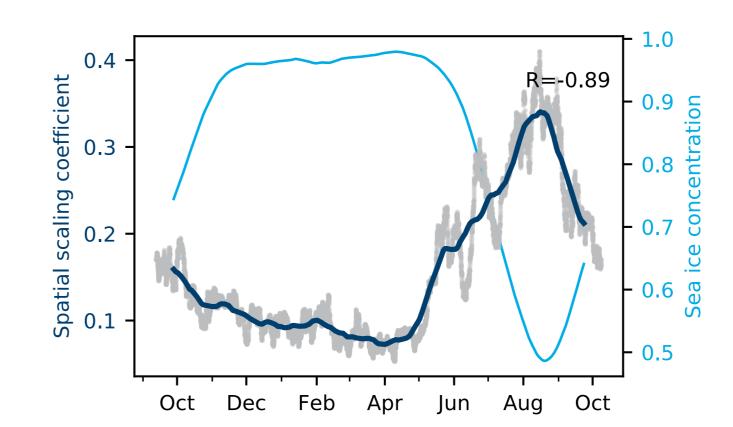


Figure: Seasonal variation of spatial scaling exponent and Arctic wide sea ice concentration

Seasonal variation

- In winter stronger stress propagation due to:
 - Higher ice strength: dense and thick ice (\checkmark modelled)
- Confinement by coastlines (\checkmark modelled)
- Stable atmospheric conditions (included in forcing)
- In summer local failure intensifies

Regional variation

• Heterogenous sea ice deformation in regions with

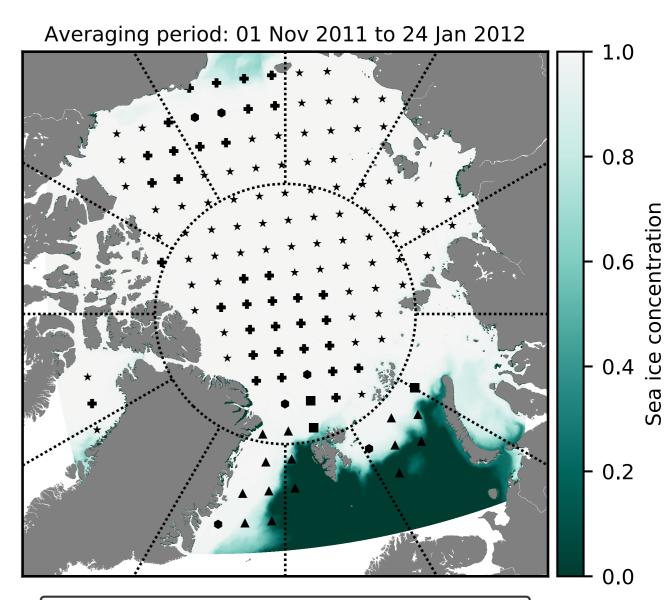
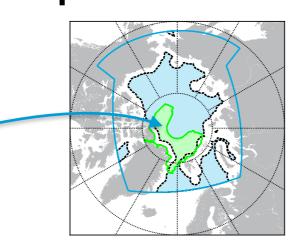


Figure: Spatial scaling properties of sea ice deformation in the model and for Envisat Geophysical Processor System (EGPS) drift data. The analysis is limited to EGPS region.

- Agreement of model results with observations regarding the amplitude of sea ice deformation and **space-time coupling**
- Different scaling exponents are influenced by
- small region with high sea ice drift (*marked in green*)
- unfiltered EGPS data → overestimation



- high sea ice drift (Fram Strait, Beaufort Sea)
- open boundaries (Barents Sea)
- Homogenous sea ice deformation in regions with
 - high ice strength (Central Arctic)
 - confinement by coasts (Laptev) Sea)

Figure (right): Regional variation of spatial scaling exponent and sea ice concentration

Spatial scaling exponent • $0.17 < \beta < 0.21$ $0 < \beta < 0.09$ $0.09 < \beta < 0.13$ • $0.21 < \beta < 0.5$ • $0.13 < \beta < 0.17$

References

Girard, L., Weiss, J., Molines, J. M., Barnier, B., & Bouillon, S. (2009). Evaluation of high-resolution sea ice models on the basis of statistical and scaling properties of Arctic sea ice drift and deformation, 114, 1–15. http://doi.org/ 10.1029/2008JC005182 Marsan, D., Stern, H., Linsdsay, R., and Weiss, J. (2004). Scale dependence and localization of the deformation of Arctic sea ice. Physical Review Letters, 93, http://doi.org/10.1103/PhysRevLett.93.178501 Rampal, P., Bouillon, S., Ólason, E., & Morlighem, M. (2016). neXtSIM: a new Lagrangian sea ice model. The Cryosphere, 10, 1055–1073. http://doi.org/10.5194/tc-10-1055-2016

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