

# The multiple roles of dust in ocean biogeochemistry

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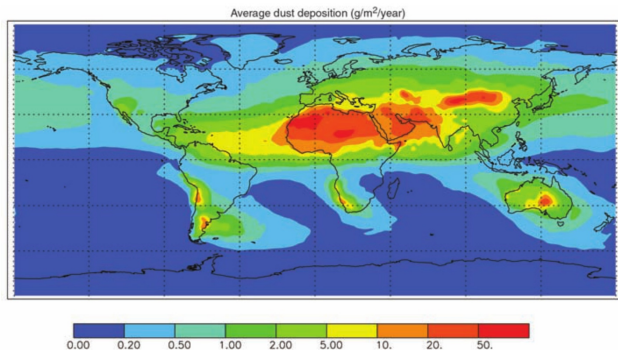
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Ocean — Climate — Sustainability  
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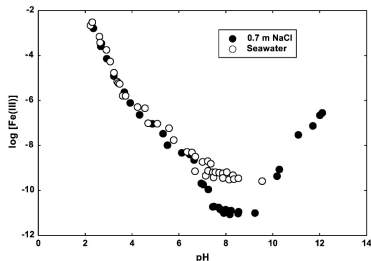
## DUST AS NUTRIENT INPUT



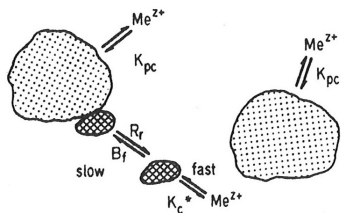
(Jickells et al., 2005)

- dust deposition brings in N (largely anthropogenic), P (lithogenic), Fe (lithogenic, combustion aerosols), other trace elements
- globally, the input is minor compared to internal recycling for elements with long residence time, such as N and P
- for elements with short residence time, such as Fe, it is important

## IRON IS PARTICLE REACTIVE



(Liu and Millero, 2002)



(Honeyman and Santschi, 1989)

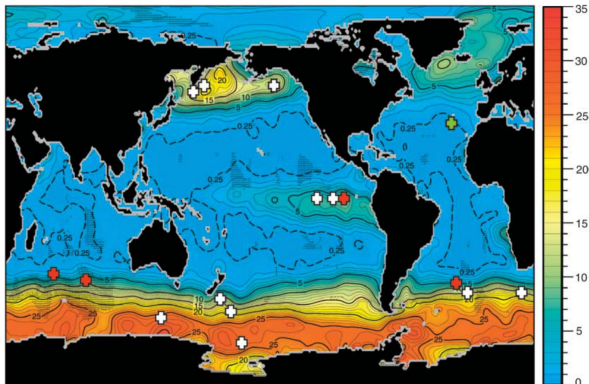
iron in oxic seawater is mostly Fe(III)

solubility of inorganic Fe(III) is extremely low:  
<  $0.1 \text{ nmol kg}^{-1}$  at typical ocean pH

→ iron is lost much faster from the ocean than nitrogen or phosphorus

loss proceeds via *scavenging* on particle surfaces, accelerated by aggregation of small colloids

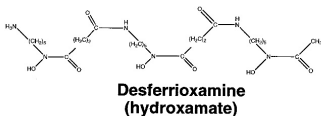
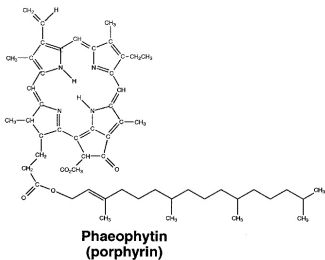
# CONSEQUENCE OF THE SHORT RESIDENCE TIME OF Fe



(Boyd et al., 2007)

- High-Nutrient-Low-Chlorophyll regions: despite enough nitrate and phosphate little net primary production
- what is missing is iron; crosses mark *iron fertilization* studies

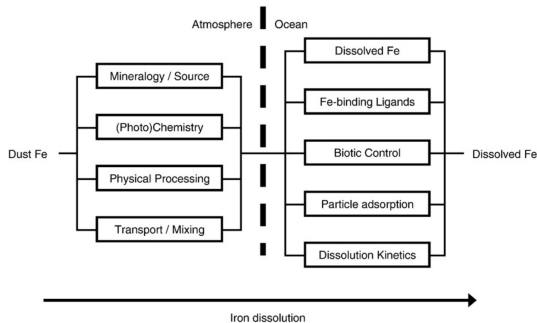
# FE-BINDING LIGANDS AS NATURE'S REMEDY



Witter et al., 2000

- typically,  $\gg 90\%$  of dissolved Fe is bound to organic ligands
- binding of iron to organics prevents rapid scavenging
- two main types of ligands: degradation products, such as porphyrins, and siderophores, produced by bacteria under iron stress
- production / degradation pathways probably as varied as ligand origins

# ATMOSPHERIC AND OCEANIC CONTROLS OF DUST INPUT

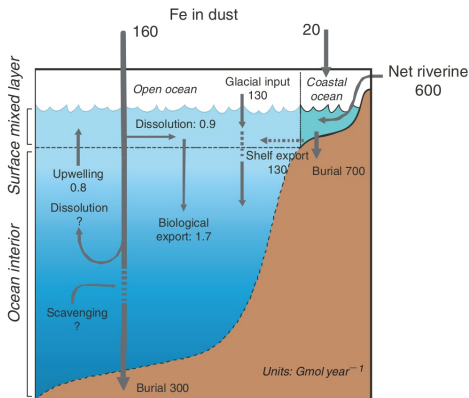


conceptual model for processes affecting solubility of dust-deposited iron (Baker and Croot, 2010)

Fraction of Fe in dust that dissolves varies from 0.1% to >10 %  
Which processes in atmosphere/ocean determine this?

→ some process understanding, but more need for exchange between atmospheric/oceanic communities!

## OTHER SOURCES OF IRON



(Hunter, 2007)

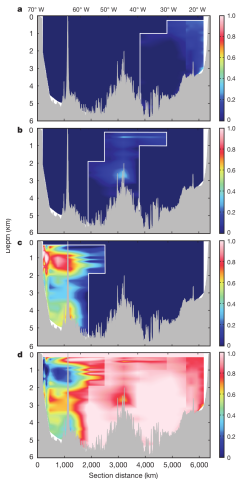
besides lithogenic dust, there are also other sources of iron:

- rivers
- sediment diagenesis
- hydrothermalism
- cosmic dust . . .

but for all of them, most of the iron is lost as particles close to the source.

quantification of different iron sources is ongoing, large uncertainties so far!

# TROPICAL ATLANTIC: DOMINATED BY DUST



(Conway et al, 2014)

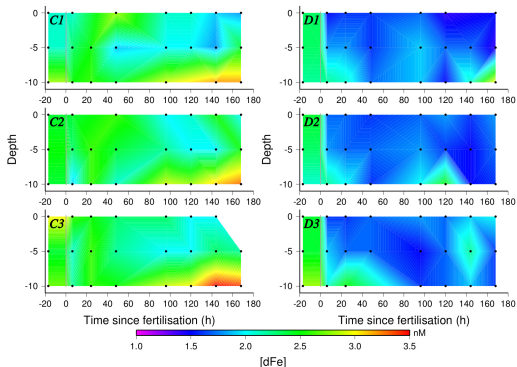
relative role of the different iron sources along a section across the subtropical/tropical Atlantic estimated from isotopic composition of dissolved iron

- sediment diagenesis
- hydrothermalism
- suspended sediment particles
- saharan dust

→ in the tropical and subtropical Atlantic, dust dominates as source of iron

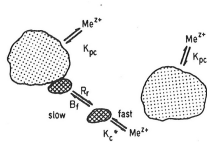


## BUT: DUST ALSO SCAVENGES DISSOLVED IRON



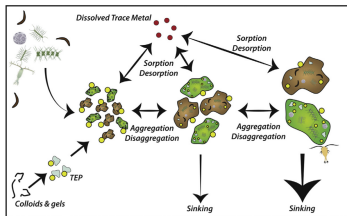
dissolved iron decreases after dust addition in mesocosms;  
dust can act as dFe sink (Wagener et al. 2010)

is that important in the open Atlantic, where often biogenic particles  
dominate?  
needs understanding & modelling of particle dynamics!

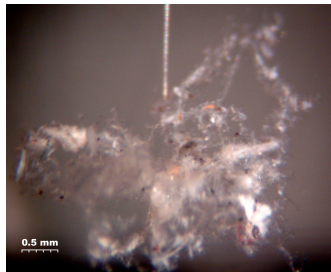


colloidal pumping  
mechanism (Honeyman  
& Santschi 1989)

## PARTICLE DYNAMICS



aggregation processes (Jackson and Burd 2015)



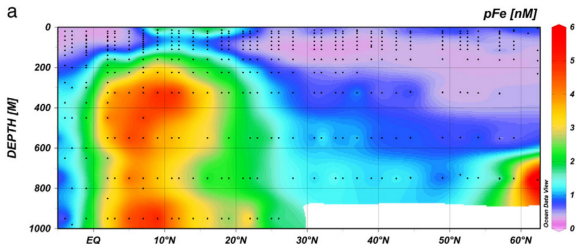
typical marine aggregate (Iversen, pers. comm.)

dust brings in mostly micrometer-sized particles

these hardly sink on their own

sinking dominated by larger, mixed organic/inorganic aggregates

# LITHOGENIC MATERIAL IN THE ATLANTIC



particulate Fe along section A16N (Barrett et al. 2012)

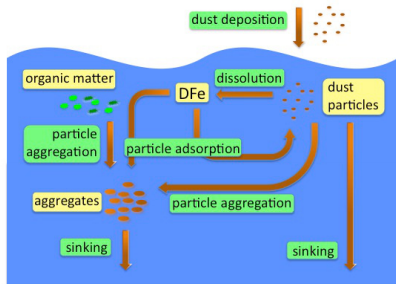
much new information on lithogenic particles from A16N and US Geotraces Atlantic Zonal Transect (Barrett et al. 2012, 2015, Lam et al. 2015, Ohnemus et al. 2015)

- increased pFe under dust plume
- high pFe at the surface, minimum around 100m depth, then again increase
- deep lithogenic particle concentration dominated by small particles
- large fraction of lithogenics highest around 100 m depth, higher towards African coast

indicates dynamic aggregation / disaggregation cycle

## MODEL SETUP

global biogeochemical model REcoM including the iron cycle (Hauck et al. 2013, Völker and Tagliabue 2015)



added model for lithogenic particles with two size classes (fine dust and faster-sinking aggregates)

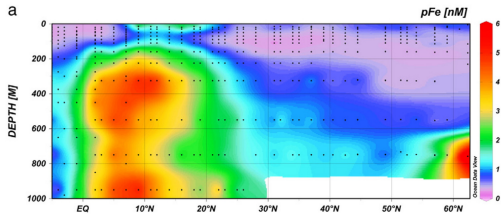
quadratic aggregation and linear disaggregation of particles

lithogenic particles included as additional scavenging agents for dissolved iron

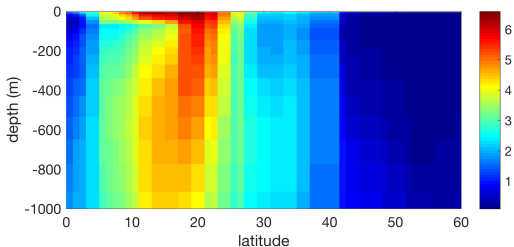
scavenging proportional to particle concentration

rate equal for organic and lithogenic particles

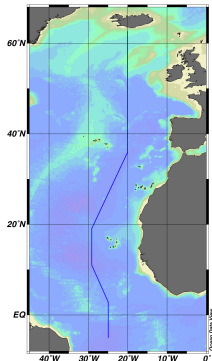
## MODELLLED VS. OBSERVED PARTICULATE Fe



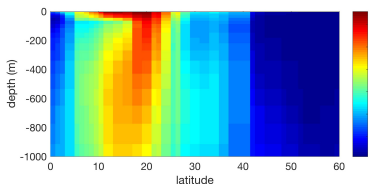
particulate Fe along section A16N (Barrett et al. 2012)



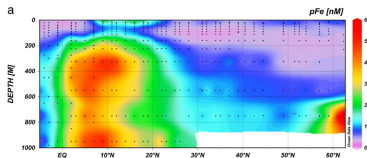
modelled particulate Fe (nM) along section A16N



# MODELLLED VS. OBSERVED pFe



modelled pFe (nM) along A16N

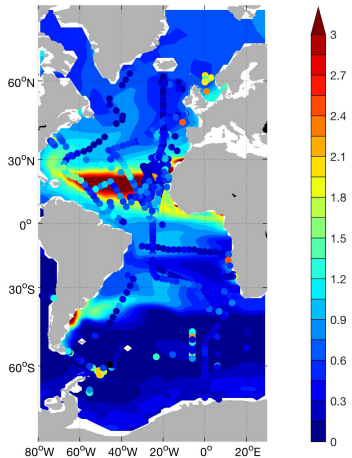


obs'd pFe along A16N (Barrett et al. 2012)

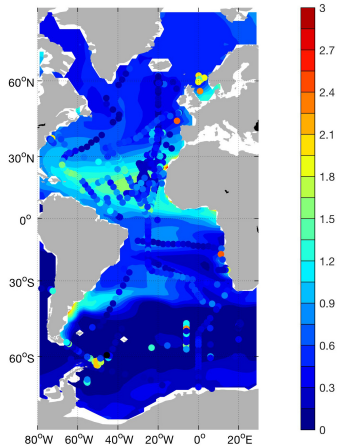
some agreement, but also some disagreement:

- + pFe concentration in the right order of magnitude
- + minimum in particle concentration around 100m depth
- + size-class distribution consistent with Ohnemus et al. (2015)
- surface pFe concentration somewhat high → aggregation rate?
- deep pFe maximum too deep → variable disaggregation?
- deep pFe maximum too far north → dust deposition?
- shelf-derived nepheloid layers absent

# EFFECT ON DISSOLVED FE

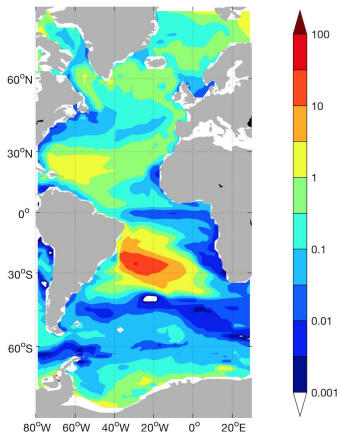


dFe with dust only as Fe source



dFe with dust as Fe source and as additional scavenging

## WHY THE REDUCTION? RESIDENCE TIME OF dFe



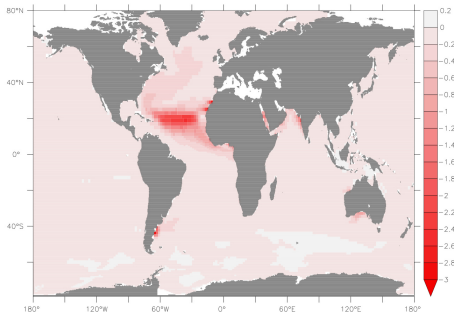
residence time (stock/total loss rate in years) of dissolved iron varies by several orders of magnitude

affected by scavenging on dust/biological particles and biological uptake

distribution of residence time agrees quite well with data-based estimates (Usher et al. 2013)



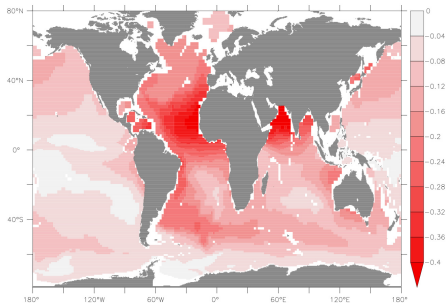
## GLOBAL EFFECT AT SURFACE



- scavenging by lithogenic particles
- reduces surface dFe directly in high-deposition regions
- but hardly everywhere else

surface dFe difference between a run with/without lithogenic particles present as scavenging agent

## GLOBAL EFFECT AT DEPTH



dFe difference (2000 m depth) between a run with/without lithogenic particles

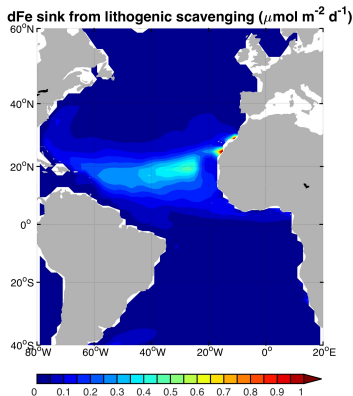
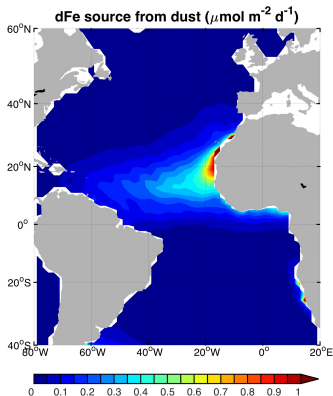
caveats: strength of effect depends on scavenging residence time, at present highly tuned in ocean iron models (Tagliabue et al., 2016) also affected by ligand excess (Völker and Tagliabue, 2015)

- dFe reduction more widespread at depth
- 1st cause: lateral transport of fine lithogenic material
- 2nd cause: downstream effect of localized scavenging
- reduces deep water dFe Atlantic – Pacific gradient

## SOME CONCLUSIONS

- dust is a major source for Fe to the ocean, especially far from land
- but relation is complicated: not only does solubility of Fe in dust vary, but dust can also scavenge iron
- lithogenic particles in the Atlantic modeled with a 2-size-class model
- both aggregation and disaggregation important, some main features of observed lithogenic particle distribution reproduced
- brings surface dFe distributions under the dust plume more in line with observations, global deep dFe distribution affected through lateral transport
- how does that affect our understanding of the role of dust in other climate states (e.g. glacial)?

# IS DUST A SOURCE OR A SINK OF dFe?



so: how much source, how much sink?

generally, dFe source stronger than vertically integrated sink; but depends somewhat where you look!